

ORGANIC-WALLED PHYTOPLANKTON BIOSTRATIGRAPHY OF THE EOCENE - OLIGOCENE TRANSITION IN THE KALLO BOREHOLE AND THE RUPELIAN STRATOTYPE AREA (NORTHWESTERN BELGIUM)

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ABSTRACT. A detailed study of the assemblages of organic-walled phytoplankton in the Eocene-Oligocene transition beds in the Kallo borehole and its comparison with the assemblages around the same transition in boreholes at Niel, Hingene-Wintham and Terhagen (the Rupelian stratotype area), reveals a hiatus between the Wintham Silt Member and the Ruisbroek Sand Member (sensu Steurbaut, 1986, 1992) in the stratotype area. In the Kallo section the sedimentation history from the latest Eocene into the Early Oligocene is more completely represented than in the stratotype area of the Rupelian.

KEYWORDS : Eocene-Oligocene transition, dinocysts, biostratigraphy, Belgium.

RESUME. Les assemblages de phytoplancton à paroi organique ont été étudiés en détail dans les dépôts de la transition Eocène-Oligocène du sondage de Kallo. La comparaison avec les sondages de Niel, Hingene-Wintham et Terhagen (région-type du Rupélien), indique l'existence d'une lacune dans la région-type entre le Silt de Wintham et les Sables de Ruisbroek (sensu Steurbaut, 1986, 1992). Dans le sondage de Kallo, la sédimentation de l'Eocène terminal jusque à l'Oligocène inférieur est plus complète que dans la région-type du Rupélien (Nord-ouest de la Belgique).

MOTS CLES : Passage Eocène-Oligocène, dinokystes, biostratigraphie, Belgique.

1. INTRODUCTION

A first rough exploration of the organic-walled phytoplankton across the Eocene-Oligocene transition in the Belgian basin was carried out by De Coninck (1986 and 1995), respectively in the Woensdrecht borehole (southern Netherlands) and in the Kallo borehole 27E-148 (near Antwerp - Belgium). From the comparison of the phytoplankton assemblages in both boreholes it was concluded (De Coninck, 1995, figure 2) that a hiatus occurs at Woensdrecht between -220 and -219 m, which corresponds approximately with deposits in the Kallo borehole between approximately -110 and -106 m. A more detailed examination of the deposits spanning the Eocene-Oligocene transition in the Kallo borehole has been undertaken. Furthermore, in the Rupelian stratotype area, about 15 to 20 km to the south of Antwerp (Figure 1), some samples were examined representing the lithostratigraphic units (Watervliet Clay Member, Wintham Silt Member and Ruisbroek Sand Member) which span the same transition in three boreholes : the Niel borehole 43W-270 (stratotype of the Ruisbroek Sand Member), the Hingene-Wintham borehole 42E-212 (stratotype of the Wintham Silt Member) and the Terhagen borehole 58W-213. These new investigations

are complementary to the detailed nannofossil biostratigraphy presented by Steurbaut (1986 and 1992) and to his sequence-stratigraphical interpretation of the Eocene-Oligocene transitional deposits in the Belgian basin (Steurbaut, 1992).

From the present study, it appears that the Kallo borehole 27E-148 is a good reference section for the Eocene-Oligocene transition in the southern North Sea basin.

2. BIOSTRATIGRAPHIC RESULTS

2.1. KALLO BOREHOLE 27E-148 (Figure 2) [all tables at end of text]

2.1.1. Samples and investigation techniques

De Coninck (1995) already examined the Eocene-Oligocene transitional deposits in the Kallo borehole in four samples only, taken at -115 m, -108 m, -98 m and -91 m. The more detailed study presented here is based on the examination of sixteen samples : -115 m, -112.1 m, -111 m, -110.5 m, -110 m, -109.5 m, -108.2 m, -108 m, -107 m, -106 m, -100 m, -98 m, -97 m, -96 m, -94 m and -92.6 m. According to Steurbaut (1992, fig. 7), this series of samples represents the top part of

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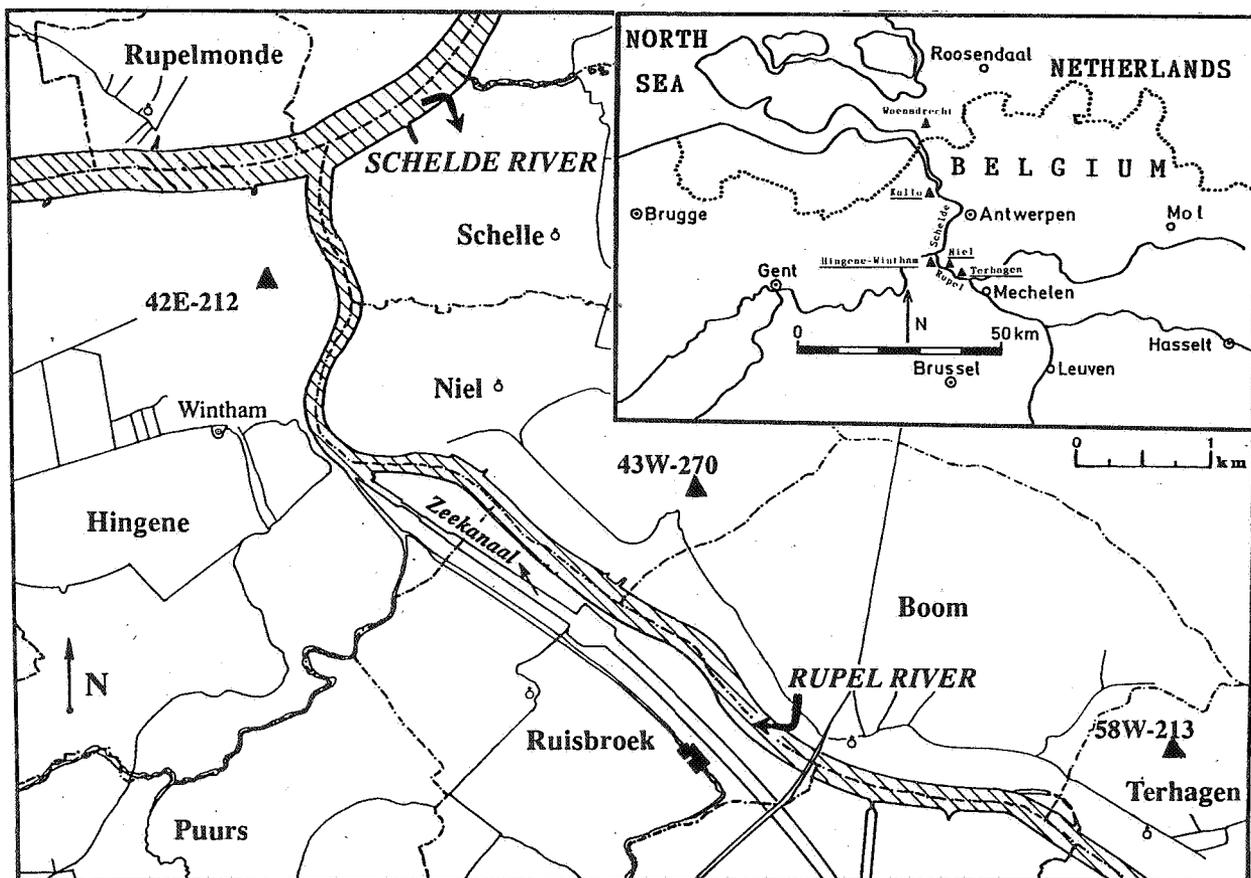


Figure 1. Localisation of the boreholes of Kallø, Hingene-Wintham, Niel and Terhagen.

the Bassevelde Sand Member (-115 m and -112.1 m), the Watervliet Clay Member (-111 m and -110.5 m), the Wintham Silt Member (-110 m?, -109.5 m), the Ruisbroek Sand Member (-108.2 m up to -96 m), the Boom Clay Formation (probably Belsele-Waas Clay Member) (-94 m and -92.6 m) (see Figure 2). The sampling gap between -106 m and -100 m could not be filled because no material was available from that part.

Table 1 presents the distribution of the Chlorophyceae, Dinophyceae, Prasinophyceae and Acritarcha in the examined section. Frequencies are indicated by the following symbols :

. sporadic <0.2 %; ; rare >0.2 % to 0.7 %; I infrequent >0.7 % to 3 %; II frequent >3 % to 10 %; X common >10 % to 25 %; XX abundant >25 %.

The letters J, C, P, Y or L in front of certain taxa indicate reworking from respectively Jurassic, Cretaceous, Palaeocene, Ypresian or Lutetian deposits. At the end of Table 1 are indicated the numbers of species reworked from these older deposits, and the numbers of species considered as having been produced during the sedimentation of the deposits.

Fluctuations of the absolute number of reworked species, which in fact show the same tendencies as the changes of the summed relative frequencies of these

taxa, give us a clue for the interpretation of changing intensities of erosion or of changing eustatic sea-levels. The occurrence at a few levels of rather frequent *Pediastrum* (fresh water planktonic green algae) can help us to decide for certain interpretations. Relative frequencies of the reworked taxa and of *Pediastrum* are also presented apart in Table 3; the fluctuations of the numbers of reworked species and high frequencies of *Pediastrum* are indicated in Figure 3.

Presence and frequency changes of plant tissue fragments in the residues are not listed, because these were not recorded systematically but they will be mentioned in the interpretation of the results.

2.1.2. Significant changes in the series of assemblages from the Kallø borehole

Among the dinocysts considered as contemporaneous with the sedimentation, a series of taxa were selected which might be useful for biostratigraphic application in the Eocene-Oligocene transitional deposits of the southern North Sea basin. Their distribution in the Kallø section is indicated in Table 2.

These selected species are used to correlate biostratigraphically the lithostratigraphical units defined by Steurbaut (1986 and 1992) in the Niel borehole 43W-

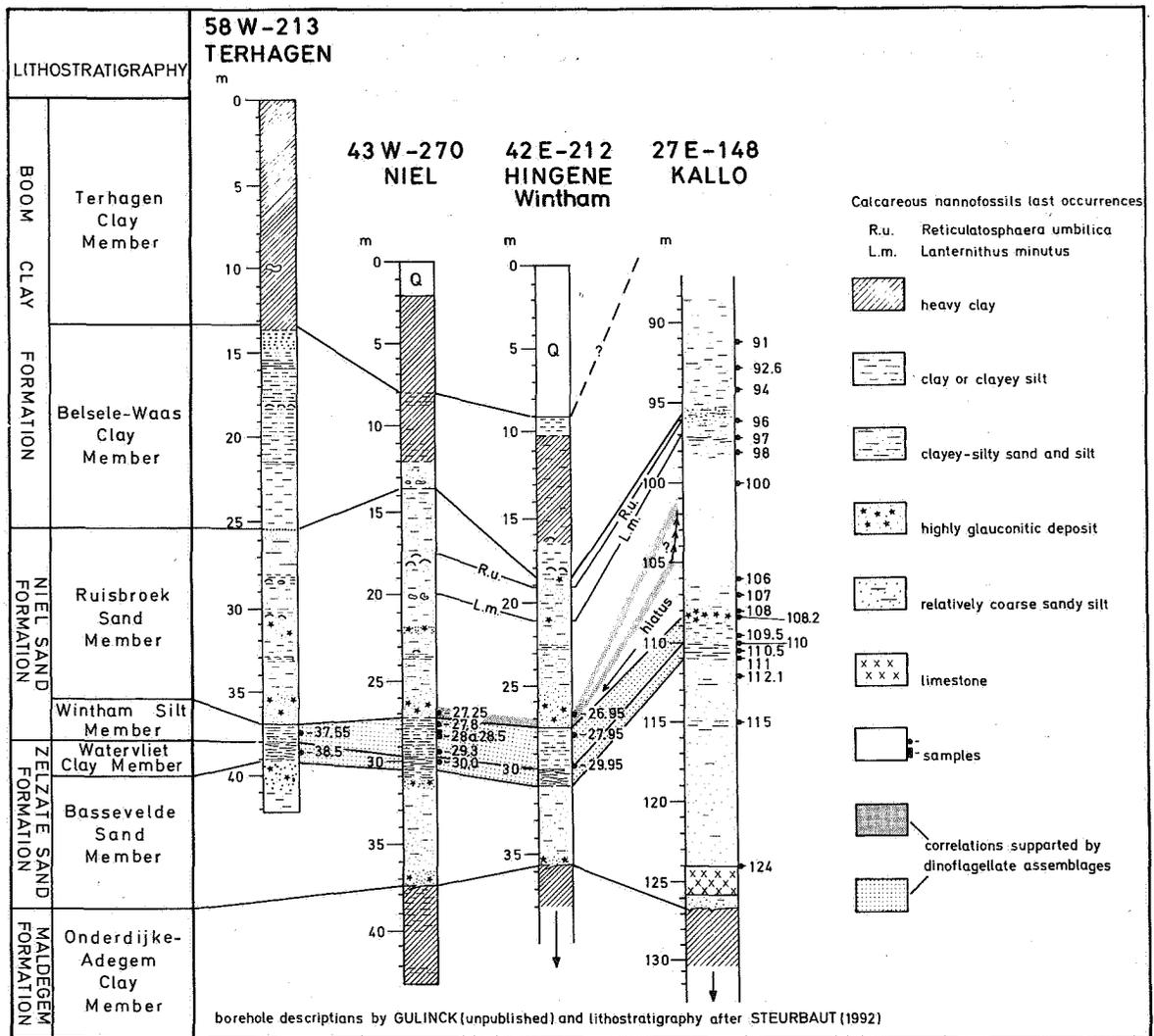


Figure 2. Lithology and lithostratigraphy of the studied borehole sections; biostratigraphic correlations.

270 and in the Hingene-Wintham borehole 42E-212 with the Kallo section (Figure 2).

Apart from these correlations, some accelerated or sudden changes in the composition of assemblages of significant taxa can reflect eventual decrease or interruptions of sedimentation, removal of deposits or changing hydrographical conditions affecting the southern North Sea basin.

When one examines Table 2, one notices a sudden change between -109.5 m and -108.2 to -108 m :

At -109.5 m are recorded the last occurrences of *Areosphaeridium diktyoplokos* (rather frequent), *Cerebrocysta bartonensis*, *Heteraulacacysta porosa*, *Membranophoridium aspinatum* forma A, *Rhombodinium perforatum*, *Homotryblium caliculum* and *Rhombodinium glabrum*. Other species such as *Charlesdowniea clathrata-coleothrypta*, *Glaphyrocysta* aff. *inculta* are still found in some higher levels but only sporadically. *Impletosphaeridium* sp. aff. *I.* sp. A - *Reticulatosphaera?* sp. A is recorded a last time at -108.2 m.

In -108.2 m to -108 m, one encounters the first *Impagidinium maculatum*, *Gerlachidium* spp. indet., *Hystriocholpoma* aff. *cinctum*, *Gerdiocysta conopeum* and *Operculodinium?* sp. cf. *O. placitum*. *Membranophoridium aspinatum* forma B and *Selenopemphix selenoides*, which were sporadically recorded at some deeper levels, are regularly found at -108.2 m and higher. The sudden change between -109.5 m and -108.2 m is also marked by a sharp decrease of reworked Mesozoic (Jurassic and Cretaceous) dinocyst taxa and by the sudden appearance of frequent *Pediastrum* (freshwater green algae) (Table 3 and Figure 3). I will try to give an interpretation of this event after having correlated the Niel-, Hingene-Wintham- and Terhagen sections with the Kallo section.

Above -108 m, the changes among the significant taxa are rather gradual up to level -97 m.

Between -97 m and -94 m, one notices an accelerated change involving the disappearance of *Hystri-*

chokolpoma aff. *cinctum*, a rarefaction of *Membranophoridium aspinatum* forma C, a temporary rarefaction of *Wetzeliella symmetrica*, *Vozzhennikovia cearaichia* and *Reticulatosphaera?* sp. A. *Membranophoridium aspinatum* forma D first appears at -97 m and disappears already above -92.6 m. *Achilleodinium* aff. *biformoides* (only once recorded at -107 m) and *Vozzhennikovia spinula* are observed in -96 m and all higher levels.

Chiropteridium spp. and *Horologinella?* sp. 1 make their first appearance at -94 m. *Apteodinium spiridioides* is recorded at -94 m only, but rather frequently. *Charlesdowniea limitata* is sporadically recorded at -94 m only.

These changes between -97 m and -94 m correspond with a certain decrease of the numbers of reworked taxa, which attained a sudden peak at -97 m (Table 3 and Figure 3). Further discussion of the changes between -97 m and -94 m is delayed, until after the section dealing with the correlation of the Niel-, Hingene-Wintham- and Terhagen deposits with the Kallo section.

2.2. NIEL BOREHOLE 43W-270 (Figure 2)

Steurbaut (1986, p. 55 and Table 2) defined the Ruisbroek Sand Member in the Niel borehole as occurring between -29.50 m and -13.20 m below the surface. At -29.50 m, a sharp boundary corresponds with the top of the Watervliet Clay Member. In 1992, Steurbaut (p. 289, 295 and figure 2) reduced the Ruisbroek Sand Member in its stratotype, the Niel borehole, to a somewhat thinner part of the section, between -27.25 m and -13.20 m: he intercalated between -29.50 m (top of the Watervliet Clay Member) and -27.25 m (base of the newly delimited Ruisbroek Sand Member) a new unit, the Wintham Silt Member, of which the stratotype was chosen in the Hingene-Wintham borehole 42E-212.

Five samples have been examined in the Niel section: -30.00 m (top part of the Watervliet Clay Member), -29.30 m, -28.5 to -28 m, -27.80 m (Wintham Silt Member) and -27.25 m (base of the Ruisbroek Sand Member).

Several of the taxa which, after the study of the Kallo section, are considered as potentially significant for biostratigraphic correlations, have been recorded in the samples from Niel. Their occurrence is presented in Table 4.

A comparison of these occurrences with the distribution in the Kallo borehole (Table 2) reveals that the groups of significant species, at -30.00 m (top part of the Watervliet Clay Member) and at -29.30 m, -28.5 to -28 m, and -27.80 m (Wintham Silt Member) which differ only slightly from each other, characterize more precisely the assemblages at -110.5 m and at -110 m in the Kallo borehole. Biostratigraphically no significant

change (e.g. hiatus or the effect of a hydrographical change) is detected between the Watervliet Clay Member and the Wintham Silt Member in the Niel borehole.

Higher in the sequence however, the difference between the assemblage at -27.80 m (top part of the Wintham Silt Member) and the one at -27.25 m (lowermost part of the Ruisbroek Sand Member) is striking: the significant species composition at -27.25 m corresponds best with the one found in the Kallo borehole at -100 m and -98 m. As no samples for examination were disponible from the Kallo borehole between -106 and -100 m, we think that the base of the Ruisbroek Sand Member at Niel can only imprecisely be correlated with a part of the Kallo section somewhere between -106 and -98 m. This means that in the Niel borehole there exists a hiatus between the Wintham Silt Member and the Ruisbroek Sand Member. This hiatus is represented at Kallo by at least 3 meters and at most 10 meters of sediments.

2.3. HINGENE-WINTHAM BOREHOLE 42E-212 (Figure 2)

Steurbaut (1992, p. 289, 295 and fig. 2) has designated this Hingene-Wintham borehole as stratotype of the Wintham Silt Member which rests with a sharp boundary at -30.00 m on the Watervliet Clay Member, and is followed at -27.55 m by the Ruisbroek Sand Member. Three samples have been examined in the Hingene-Wintham borehole: -29.95 m (lowermost part of the Wintham Silt Member) and -27.95 m (top part of the Wintham Silt Member), and -26.95 m (lower part of the Ruisbroek Sand Member). The biostratigraphically significant taxa in these three samples are indicated in Table 5.

At -29.95 m and -27.95 m (Wintham Silt Member) the assemblages of significant species suggest a correlation with -111 to -109.5 m in the Kallo borehole. A similar species composition has also been encountered in the top of the Watervliet Clay Member and in the Wintham Silt Member of the Niel borehole 43W-270 (see higher). The assemblage at -26.95 m, from the Ruisbroek Sand Member some 60 cm above its base, presents several significant species: *Glaphyrocysta semitecta*, *Wetzeliella symmetrica*, *Membranophoridium aspinatum* forma C and a form which seems intermediary between our *Svalbardella* sp. indet. and *Palaeocystodinium golzowense* are regularly encountered. *Vozzkennikovia cearaichia* and *Reticulatosphaera?* sp. A appear only sporadically. Only a few *Achilleodinium* aff. *biformoides* and *Chiropteridium* sp. indet. are recovered. Apart from the two last species, no *Membranosphaeridium aspinatum* forma D nor other significant species which appear at Kallo -97 m or higher are found in the assemblage.

Most of the significant species mentioned can be found together at Kallo from -106 to -98 m. The relatively frequent *Svalbardella-Palaeocystodinium*-like specimens suggest a correlation with the lower part of this interval. Only the presence of some *Chiropteridium* specimens, never recorded at these levels at Kallo, where this form is found a first time at -94 m, is confusing. The Wintham-Hingene assemblage at -26.95 m might eventually give evidence of an earlier short term appearance of *Chiropteridium*, which in that case was not observed in the Kallo trajectory because of sampling gaps.

A correlation of -26.95 m with some level at Kallo between -106 and -100 m, rather close to -106 m, seems nevertheless the best option, especially when one considers also the data from lithostratigraphic and calcareous nannoplankton studies by Steurbaut (1986, table 2, 1999, fig. 2 and p. 297). In our Figure 2 the last occurrences of the nannofossil species *Lanternithus minutus* and *Reticulofenestra umbilica* are indicated in the three boreholes.

The proposed correlation corresponds approximately with the one concluded for the base of the Ruisbroek Sand Member in the Niel borehole. If correct, it reveals again a certain hiatus between the Wintham Silt Member and the Ruisbroek Sand Member, apparently a marked break of continuity in the type area of the Rupelian.

2.4. TERHAGEN BOREHOLE 58W-213 (Figure 2)

Steurbaut (1986, p. 58) had designated this Terhagen borehole as a parastratotype for the Ruisbroek Sand Member. After having introduced the Wintham Silt Member (Steurbaut, 1992, p. 295), the succession of lithostratigraphic units in the Terhagen borehole was interpreted by Steurbaut (*ibid.*, fig. 2) as follows: the boundary between the Watervliet Clay Member and the Wintham Silt Member is situated at about -37.80 m, the boundary between the Wintham Silt Member and the overlying Ruisbroek Sand Member at -37.0 m.

Organic-walled phytoplankton has been examined in two samples: -38.50 m in the middle of the Watervliet Clay Member, and -37.55 m in the middle of the Wintham Silt Member. The biostratigraphically significant species recorded in both samples are indicated in Table 6.

The significant species association recorded at -38.50 m is found in the Kallo borehole between -115 and -111 m. The association at -37.55 m corresponds best with these from -111 m to -109.5 m in the Kallo borehole. Only little differences exist between both associations of significant species, indicating once more that between the Watervliet Clay Member and the

Wintham Silt Member no biostratigraphical hiatus can be detected.

3. TENTATIVE INTERPRETATION OF THE KALLO BOREHOLE; CONFRONTATION WITH THE OBSERVATIONS IN THE OTHER BOREHOLES

The data obtained from the boreholes Niel, Hingene-Wintham and Terhagen are valuable for the interpretation of the changes observed in the Kallo borehole (Table 3 and Figure 3).

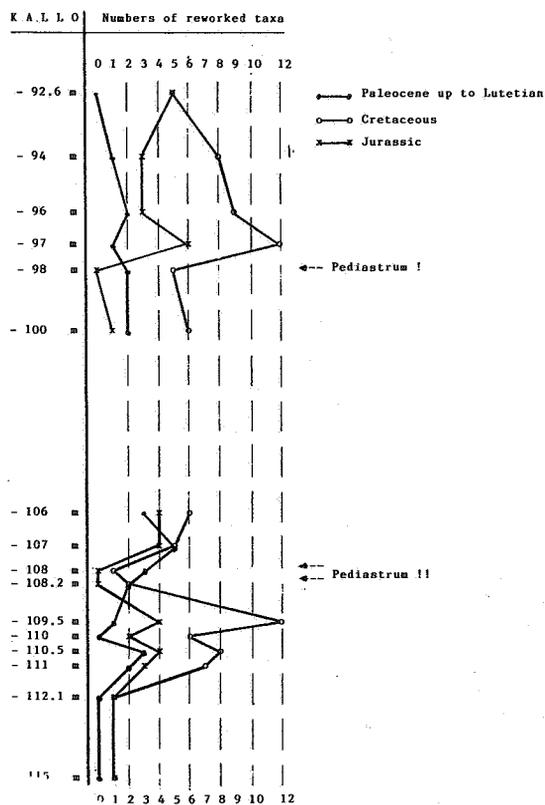


Figure 3. Numbers of reworked dinocyst taxa and frequencies of *Pediatrum* in the Kallo section.

-115 up to -109.5 m

There are no significant changes in the composition of the contemporaneous organic-walled phytoplankton assemblages from the top part of the Bassevelde Sand Member over the Watervliet Clay Member to the top of the Wintham Silt Member. No biostratigraphic hiatuses can be detected between these units. When one examines the assemblages of the Watervliet Clay and Wintham Silt Members in the Niel- and Terhagen boreholes, one can not differentiate them on their contents of contemporaneous phytoplankton. The assem-

blages contain the same significant species as those in Kallo between -115 and -109.5 m.

Unlike the monotony of the composition of significant species, one notices in this part of the Kallo traject a marked contrast between the low frequencies of reworked phytoplankton at -115 m and -112.1 m, on the one hand, and their high frequencies, especially of Mesozoic taxa, from -111 m up to -109.5 m, on the other hand. Between both parts of the traject, I observed also the absence of plant tissue fragments in -115 m and -112.1 m, followed by their presence, although not frequent, in -111 m up to -109.5 m. These differences can be explained by a suddenly lowered sea level. The reworked taxa (as well more diverse as more frequent) offer evidence of intensified erosion, while the presence of plant tissue fragments indicates the nearness of the coast during the sedimentation of the Watervliet Clay and Wintham Silt Members.

In the preceding period, when the top part of the Bassevelde Sand Member was deposited (-115 m and -112.1 m), the sea level was probably still high, erosion of older deposits negligible and the coast far away as marine sediments of this age are present at Oret, Entre-Sambre-et-Meuse, about 100 km to the south of the Rupel type-area. The lower sea level which followed (-111 m up to -109.5 m), enhanced erosion probably of coastal cliffs to the west or north-west of the Belgian basin. The eroded material was probably carried towards the Kallo area by currents, which in the North Sea basin are normally counterclockwise. Maybe the regression was not the only cause of the erosion of unweathered deposits in which microfossils were well preserved. Ziegler (1990, p. 150 and 161) summarises some studies revealing a Late Eocene-Oligocene phase of regional uplift which affected, among others, the region of Yorkshire, its eastern prolongation towards the Sole Pit basin, also the Channel, the Celtic Sea and other parts of England. We do not know if that phase is coeval with any part of the Eocene-Oligocene transitional deposits studied here. Anyway, a cumulative effect of an eventual regional uplift to the west or north-west of the basin and the lowered sea level could explain more easily the erosion of Mesozoic rocks from which the reworked microfossils came.

In the Watervliet Clay and Wintham Silt of the Rupel type-area reworked dinocysts are less frequent and less diverse than in the corresponding deposits at Kallo. No doubt, the distribution of reworked microfossils in the Belgian basin was confined to the course followed by the current carrying eroded material.

Transition from -109.5 m to -108.2 and -108 m

This transition is marked by a prompt substitution of significant taxa (see higher) probably reflecting a tem-

porary break or strong decrease of the sedimentation. Furthermore, the number of reworked Mesozoic dinocysts declines drastically and fresh-water green algae *Pediastrum* appear suddenly and frequently at -108.2 and -108 m while plant-tissue fragments become more frequent.

The reduction or break of sediment supply can reflect a temporary rise of the sea-level relative to the nearby coastal regions. If so, drainage of the coastal plain became first of all less effective and freshwater lakes may have formed. When the further rising sea finally invaded the coastal plain area, plant tissues and *Pediastrum* from the lakes and from their bottom sediments were redistributed on the sea bottom.

In the Niel- and Hingene-Wintham boreholes, no deposits corresponding to this supposed sea-level rise are found; instead a hiatus is detected between the Wintham Silt Member and the Ruisbroek Sand Member, probably not an erosional one but rather one marking the absence of sedimentation. If we maintain the proposed model of temporarily rising sea-level and regional palaeogeographic evolution based on the observations in the Kallo section between -109.5 m and -108.2 to -108 m, the sea-level rise affected only a restricted area and was thus relatively limited. The sudden fall of the number of reworked species between -109.5 m and -108.2 m (before their renewed but less pronounced increase at -107 m and -106 m) may be explained, not only by the slight sea-level rise but also by a change in position of the currents carrying the reworked material. The discontinuity at this transition is much more important than the earlier or later changes observed in this study of the Late Eocene-Early Oligocene deposits. It corresponds furthermore with the disappearance of *Areosphaeridium diktyoplokus*, a widely observed biostratigraphic event which is a reliable candidate for the determination of the Eocene-Oligocene boundary.

-107 m and -106 m

After the former particular intermezzo, sedimentation went on in the Kallo area only, because no time-equivalent deposits are found in the Rupel type-area. A number of reworked taxa are again recorded in -107 and -106 m, while *Pediastrum* disappears and plant tissue fragments become less frequent. Maybe currents with reworked microfossils affected again more or less the Kallo area. The presence of *Svalbardella* sp. indet. in -106 m only, is worth mentioning as the genus is characteristic of relatively cold water, according to Head & Norris (1989, p. 533).

-106 m to -100 m

As no samples could be studied between -106 m and -100 m, data are lacking. In the Rupelian type-area sedimentation probably started again at some time in the course of deposition of this interval.

-100 m to -98 m

Between both levels, the number of reworked Mesozoic taxa diminishes only slightly and at -98 m *Pediastrum*, only sporadically found at -100 m, becomes relatively frequent. These moderate changes reflect probably a slight sea-level rise and a redistribution of some invaded coastal plain fresh-water deposits.

Transition from -98 m to -97 m

At -97 m we note an impressive increase of Jurassic and Cretaceous microfossils in the assemblage. The sudden arrival of this reworked material is probably the result of a regression. In the Ruisbroek Sand Member studied in boreholes near the mouth of the Rupel river, Steurbaut (1992, figs. 3 and 4) found traces of a regression marking the transition of subsequence A₁ to subsequence A₂. This regression may have been important enough to explain the increase of reworked forms. The top of subsequence A₁ is, according to Steurbaut (ibid., p. 297 and Fig. 3), characterized by the last occurrence of the calcareous nannoplankton species *Lanternithus minutus*, which was indeed found at -97 m (see Steurbaut, 1986, Table 1).

-96 to -94 m

At -96 m the number of reworked forms decreases. It possibly reflects a rising sea-level, eventually corresponding to the flooding surface 2 in subsequence A₂, according to the sequence-stratigraphical interpretation by Steurbaut (1992, fig. 7).

At -94 m appear four species: *Apteodinium spiridoides*, *Charlesdowniea limitata*, *Chiropteridium* spp. indet. and *Horologinella?* sp. 1. The frequency of *Vozzhenikovia spinula*, which was first recorded sporadically at -96 m, increases clearly at -94 m.

The change observed between -96 and -94 m corresponds in fact with the transition at -95.5 m of the Ruisbroek Sand Member (Niel Sand Formation) to the Belsele-Waas Clay Member (Boom Clay Formation). It follows here immediately the last occurrence of the calcareous nannoplankton species *Reticulofenestra umbilica* (see Steurbaut, 1986, Table 1).

-94 to -91 m

There are no significant changes to mention among the succeeding assemblages in the Belsele-Waas Clay Member.

4. CONCLUSION

From the distribution of dinoflagellate cyst species produced by the plankton at the time of sedimentation of the Bassevelde Sand Member, the Watervliet Clay Member and the Wintham Silt Member, it is obvious that biostratigraphically there are no significant hiatus between these three lithostratigraphic units.

The sharp lithological boundary between the Watervliet Clay Member and the Wintham Silt Member in the Niel, Hingene-Wintham and Terhagen boreholes, with locally deep bioturbations descending from the Wintham Silt in the Watervliet Clay, does not represent a long-time interruption of sedimentation between both units.

In the Watervliet Clay and Wintham Silt Formations the appearance of plant tissue fragments and especially at Kallo of many reworked dinocysts in the assemblages indicates a sea-level drop and the resulting approach of the coast; this in contrast with the earlier high sea-level and remote coast during sedimentation of the Bassevelde Sand Member. A major change of the assemblages, with among others the disappearance of *Areosphaeridium diktyoplokus*, is noted at the transition from the Wintham Silt Member to the overlying Ruisbroek Sand Member. Probably sedimentation stopped for a while or became very limited in the Kallo area as a result of a slight rise of the sea-level which must have hindered temporarily drainage of the nearby coastal plain. Indeed, subsequent invasion by the sea redistributed freshwater deposits with their *Pediastrum* and plant tissue fragments. At the mean time the Rupel type-area was temporarily excluded from further sedimentation. During this interval of slowly rising sea-level and restricted sedimentation, about four or a little more meters of deposits were laid down in the Kallo area where some colder water influx seems to have occurred, indicated by the presence of *Svalbardella* at Kallo -106 m. Soon afterwards, the sedimentation of the Ruisbroek Sand Member spread into the Rupel type-area. There nearly the same lithological changes and calcareous nannofossil successions are observed as in the remaining Ruisbroek Sands at Kallo. At Kallo -97 m the sudden increase of the number of reworked dinocysts in the assemblage indicates a regression corresponding with the transition of subsequence A₁ to subsequence A₂. At Kallo -96 m the decreasing number of reworked dinocysts may indicate a transgression which eventually corresponds with the flooding surface 2 in subsequence A₂. The boundary between the Ruisbroek Sand Member and the Belsele-Waas Clay Member is situated at Kallo -95.5 m and accompanied by a relatively marked change of the assemblages, although less pronounced than the change at the boundary between the Wintham Silt Member and the Ruisbroek Sand Member, which reflects the most important discontinuity in sedimentation - underlined by a hiatus in the Rupelian stratotype area - and in succession of dinocyst-assemblages in the Late Eocene-Early Oligocene sequences of North-western Belgium.

5. REMARKS ON CERTAIN TAXA NOT YET COMMENTED ON IN EARLIER STUDIES BY THE AUTHOR (DE CONINCK, 1986 AND 1995)

Achilleodinium aff. *biformoides* (EISENACK, 1954)
Pl. 7, figs. 4 and 5

Remark : Our specimens differ from *A. biformoides* (EISENACK, 1954, p. 68; pl. 11, figs. 16-20) by their elongate shape.

Dimensions of the cyst body : 70 to 78 μm long, 45 to 50 μm wide.

Elytrocysta breva STOVER & HARDENBOL 1994
Pl. 3, figs. 12, 13, 14 and 15.

Remark : *E. breva* STOVER & HARDENBOL, 1994 (p. 35; pl. 2, figs. 10-12) corresponds with *Elytrocysta?* sp. B in De Coninck, 1995 (p. 85; pl. 1, figs. 10-11) and is larger than *Elytrocysta?* sp. A in De Coninck, 1986 (p. 12; pl. 2, figs. 15-18).

Dimensions : between 37 and 45 μm .

Gerlachidium? spp. indet.
Pl. 2, figs. 6, 19 and 20.

Remark : Our *Gerlachidium?* spp. indet. bear small, short and capitate processes, some of which are relatively wide and flattened, especially these on the anterior and posterior margin of the paracingulum. The cysts general form is rather similar to that of *Lejeunecysta* spp. than to that of *G. aechmophorum* (BENEDEK, 1972).

Dimensions of the cyst : 40 à 50 μm (two specimens), 60 à 65 μm (two other specimens).

Horologinella? sp. 1
Pl. 2, figs. 11, 16 and 17.

Remark : Our *Horologinella?* sp. 1 are thin walled organisms, with a general form which resembles a quadrangular pillow which is somewhat assymetrically invaginated at one of its sides. An attribution of this species to the genus *Horologinella* COOKSON & EISENACK, 1962 (p. 271) emend. STOVER & EVITT, 1978 (p. 53) emend. BACKHOUSE, 1988 (p. 90) is questionable.

Dimension : 15 to 18 μm .

Hystrichokolpoma aff. *cinctum* KLUMPP, 1953
Pl. 4, figs. 1, 2 and 3.

Remark : Our *H. aff. cinctum* differ from the species (KLUMPP, 1953, p. 389; pl. 17, figs. 3-5a-d) by their distally widening, more tubiform processes.

Dimensions :

cyst body : \pm 50 μm

antapical process : about 30 to 40 μm long, 8 to 13 μm wide below its distal evasion.

pre- and post cingular processes : about 20 to 25 μm long, 10 to 15 μm wide below their distal evasion.

paracingular and parasulcal processes : about 15 to 20 μm long, 2 to 3 μm wide below their distal evasion.

Hystrichokolpoma aff. *rigaudae* DEFLANDRE & COOKSON, 1955

Pl. 1, figs. 15 and 16.

Remark : Our *H. aff. rigaudae* have an elongate cyst body, as our *H. cf. rigaudae* (in De Coninck, 1986, p. 13; pl. 4, fig. 5). The paracingular and parasulcal processes are relatively narrow, but not as tenuous as in *H. cf. rigaudae*.

Dimensions :

cyst body : 42 to 50 μm long; 30 to 35 μm wide.

antapical process : 25 to 30 μm long; 6 to 8 μm wide below its distal evasion.

pre- and postcingular processes : 15 to 22 μm long; 5 to 10 μm wide below their distal evasion.

paracingular and parasulcal processes : about 1 to 2 μm wide below their distal evasion.

Impletosphaeridium sp. aff. *I. sp. A* in De Coninck, 1986 -

Reticulatosphaera? sp. A in De Coninck, 1995
Pl. 2, fig. 4.

Remark : This form differs from our *Reticulatosphaera?* sp. A in De Coninck, 1995 (p. 87; pl. 6, figs. 11-16) by the shorter and thinner filaments spreading from the distally evaded extremity of the processes. It is not clear whether the filaments from neighbouring processes are connected with each other or not. In *Reticulatosphaera?* sp. A the filaments are stronger and connections are very clearly observed. The form may be considered as a precursor of *Reticulatosphaera?* sp. A and seems related to *Impletosphaeridium* sp. A in De Coninck, 1986 (p. 15; pl. 5, figs. 16-19).

Dimensions :

cyst body : about 25 to 30 μm .

processes : about 12 to 18 μm long, about 1 to 4 μm wide in the middle and 5 to 7 μm wide distally.

Litosphaeridium? aff. *mamellatum* DE CONINCK, 1977

Pl. 1, figs. 8, 9 and 10.

Remark : *L.? aff. mamellatum* is considerably larger than the species described by De Coninck, 1977 (p. 42; pl. 6, figs. 8-13). Its mamella-like processes are relatively flattened and enlarged. The wall of the central cyst body is often invisible.

Dimensions :

cyst body : about 33 to 35 μm .

processes : 5 to 8 μm high, 20 to 26 μm wide.

Overall dimensions : about 45 to 55 μm .

Membranophoridium aspinatum GERLACH, 1961

The species was first described by Gerlach, 1961 (p. 199; pl. 29, figs. 7, 8) in Oligocene deposits of Northern Germany. In the Latest Eocene and Early Oligocene deposits of the Belgian basin the morphology of *M. aspinatum* changes more or less gradually with the age of the deposits. These changes can be expressed as a series of four forms :

forma A : Pl. 1, figs. 1, 2, 3 and 4.

The central part of the dorsal and ventral surfaces of the cyst presents irregular thickenings of the periphragm. These thickenings are more frequent on the ventral surface than dorsally and compose in several specimens a kind of reticulum; dorsally an alignment in paracingular position, of these thickenings, is observed in some specimens.

forma B : Pl. 3, figs. 1 up to 11.

Irregularities of the wall thickness, which are seen in the central part of the dorsal and ventral cyst surfaces of *forma A*, have disappeared in *forma B*.

The antapical extremity of each lateral ventral flange, especially of the left one, is somewhat more protruding than the rest of the flanges. In some specimens the antapical extremity of both dorsal and ventral lateral flanges is rather conical; the left ventral one again most protruding.

forma C : Pl. 6, figs. 1 up to 5.

Compared with *forma B*, the antapical protrusion of both lateral ventral flanges is larger more like a blunt process. Especially the left one of these ventral protrusions is well developed. In some specimens this left antapical process is only in its lower part in connection with the lateral flange, and has than a rather fluted appearance.

forma D : Pl. 6, figs. 6 up to 10. Pl. 7, fig. 8.

The distal border of the lateral flanges presents several concave incisions, more frequent on the ventral flanges than on the dorsal ones. The incisions on the posterior part of the ventral flanges are in general wide and deep. Antapically, the lateral dorsal and ventral flanges are elongated into a distally closed and truncated process; ventrally, these processes (especially the left one) are more developed than dorsally and may present some rudimentary ramifications. The processes communicate only proximally with the corresponding lateral flange. *Forma D* approaches *Chiropteridium*, in which the flanges are more deeply incised and for their greater part replaced by series of hollow processes, eventually communicating with each other up to a certain height. Morphotypes A and B of «*Chiropteridium aspinatum*», described by Schindler, 1992 (p. 201-204, figs. 1, 2, 3,

4, 5) seem to correspond with *forma D*.

Selenopemphix? sp. 3 in Head & Norris, 1989
Pl. 7, figs. 11, 12 and 13.

Remark : This species was only figured by Head & Norris, 1989, pl. 7; figs. 4 and 8. They recorded it in the Early Oligocene NP22 Zone and basal part of the NP23 Zone in the Labrador Sea borehole 647A.

Tectatodinium? sp. cf. *Filisphaera* sp. A in Head, Norris & Mudie, 1989
Pl. 1, fig. 14.

Remark : This species, resembling *Tectatodinium pellitum* WALL 1967 (p. 113; pl. 16, figs. 11-12), presents a granular periphragm of slightly uneven thickness, appressed to a thin, hyaline endophragm. In some specimens the apex and the margins of a paracingulum are very weakly marked by slightly thicker periphragm. A relation with *Filisphaera* sp. A in Head, Norris & Mudie, 1989 (p. 435; pl. 7, fig. 1, 7, 11) seems likely.

Dimensions :
cyst body : about 50 to 62,5 μm .
cyst wall : about 1 to 2,5 μm thick.

Xenicodinium? sp. indet.
Pl. 1, figs. 11, 17 and 18.

Remark : Solid, conical to blunt processes, at a distance of about 1 to 3 μm from each other, are scattered on the surface of a globular cyst. The cyst wall is two-layered, with a thin endophragm and an appressed thicker periphragm. In one specimen a precingular opercular plate (3") is seen inside the cyst body; the archaeopyle seems however larger than that one opercular plate. Other specimens are often sliced. I have attributed this form tentatively to the genus *Xenicodinium* KLEMENT, 1960 (p. 53).

Dimensions :
cyst : 40 to 55 μm
endophragm : less than 1 μm thick.
periphragm : 1 to 2 μm thick.
processes : ± 1 μm long and 1 μm wide.

6. ACKNOWLEDGEMENTS

I thank Dr. P. Laga, Geological Survey of Belgium, for authorizing the study of samples from the Kallo, Niel, Hingene-Wintham and Terhagen boreholes. Special thanks are due to Dr. E. Steurbaut, at the Royal Institute of Natural Sciences of Belgium, for his stimulating cooperation and encouraging interest in the results of the study when in progress. I appreciate very much also the work and goodwill of Mr. Daniel Bavay who drew the figures and helped to assemble the plates, and of Mrs Nelly Reynaert who prepared the manuscript.

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PLATE 1

- Figures 1, 2. *Membranophoridium aspinatum* GERLACH, 1961 forma A
Kallo -110 m; slide 1. 500x.
- Figures 3, 4. *Membranophoridium aspinatum* GERLACH, 1961 forma A
Kallo -109.5 m; slide 1. 500x.
- Figures 5, 6. *Glaphyrocysta* aff. *inculta* (MORGENROTH, 1966)
Kallo -115 m; slide 5. 500x.
- Figure 7. *Rhombodinium perforatum* (JAN DU CHENE and CHATEAUNEUF, 1975)
Kallo -115 m; slide 4. 500x.
- Figures 8, 9. *Litosphaeridium* ? aff. *mamellatum* DE CONINCK, 1977
Kallo -108 m; slide 1. 500x.
- Fig. 10. *Litosphaeridium* ? aff. *mamellatum* DE CONINCK, 1977
Kallo -108 m; slide 2. 500x.
- Figure 11. *Xenicodinium* ? sp. indet.
Kallo -106 m; slide 1. 500x.
- Figure 12. *Rhombodinium draco* GOCHT, 1955
Kallo -111 m; slide 2. 500x.
- Figure 13. *Paucilobimorpha spinosa* (COOKSON, 1965)
Kallo -110 m; slide 3. 500x.
- Figure 14. *Tectatodinium* ? sp. cf. *Filisphaera* sp. A in HAED, NORRIS and MUDIE, 1989
Kallo -92.6 m; slide 2. 500x.
- Figure 15. *Hystrihokolpoma* aff. *rigaudae* DEFLANDRE and COOKSON, 1955
Kallo -106 m; slide 1. 500x.
- Figure 16. *Hystrihokolpoma* aff. *rigaudae* DEFLANDRE and COOKSON, 1955
Kallo -107 m; slide 1. 500x.
- Figures 17, 18. *Xenicodinium* ? sp. indet.
Kallo -92.6 m; slide 3. 500x.

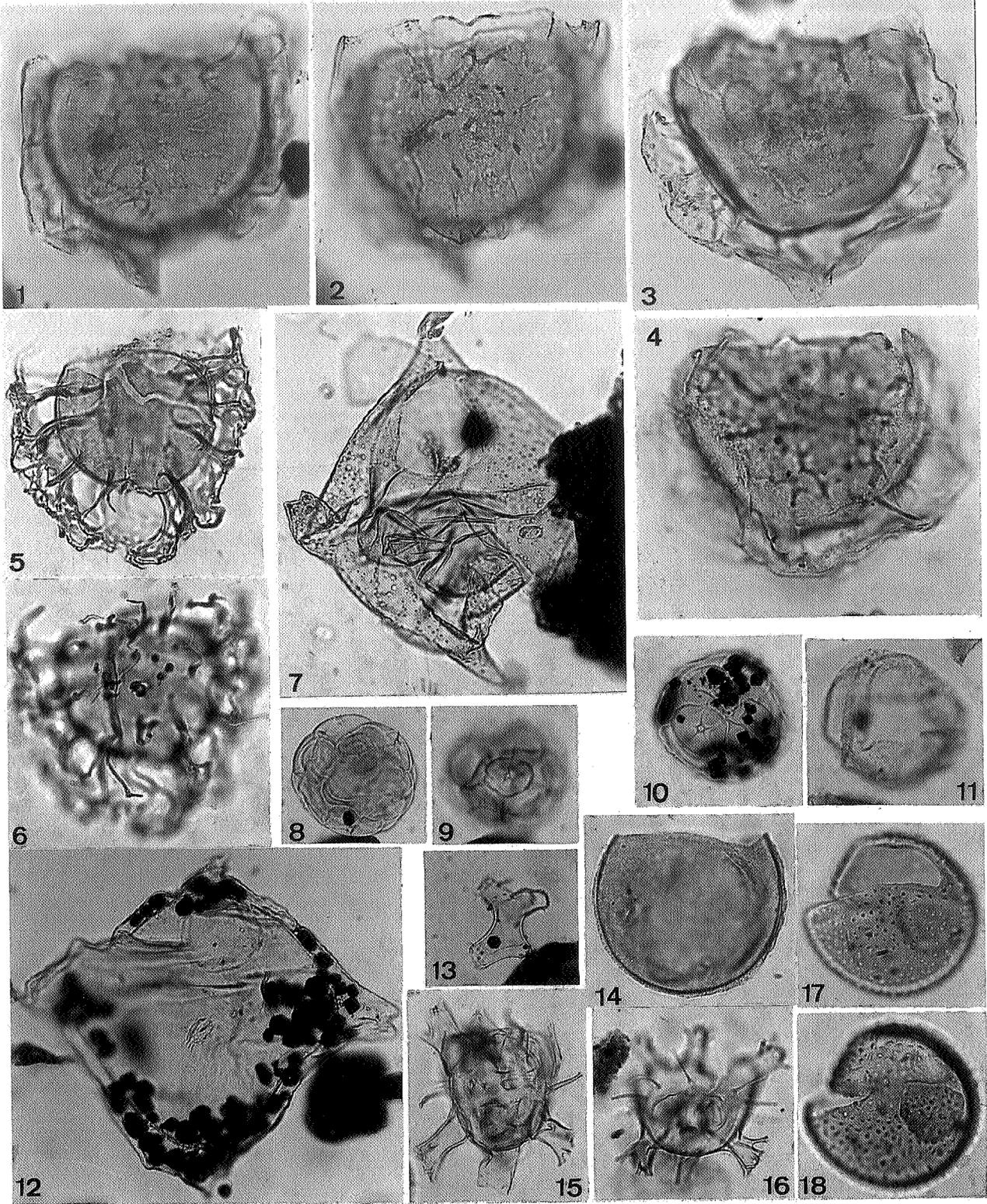


PLATE 2

- Figure 1. *Thalassiphora fenestrata* LIENGJARERN, COSTA and DOWNIE, 1980
Kallo -108.2 m; slide 1. 500x.
- Figure 2. *Thalassiphora reticulata* MORGENROTH, 1966
Kallo -110.5 m; slide 1. 500x.
- Figure 3. *Svalbardella* sp. indet.
Kallo -106 m; slide 1. 500x.
- Figure 4. *Impletosphaeridium* sp. aff. *I.* sp. A in DE CONINCK, 1986 - *Reticulosphaera* ? sp. A in DE CONINCK, 1995
Kallo -109.5 m; slide 2. 1000x.
- Figure 5. *Planoperidinium gracile* DE CONINCK, 1986
Kallo -108.2 m; slide 2. 500x.
- Figure 6. *Gerlachidium* ? sp. indet.
Kallo -100 m; slide 2. 500x.
- Figures 7, 8. *Impagidinium maculatum* (COOKSON and EISENACK, 1961)
Kallo -107 m; slide 2. 500x.
- Figure 9. *Svalbardella* sp. indet.
Kallo -106 m; slide 2. 500x.
- Figure 10. *Svalbardella* sp. indet.
Kallo -106 m; slide 1. 500x.
- Figure 11. *Horologinella* ? sp. 1
Kallo -92.6 m; slide 1. 1000x.
- Figures 12, 13. *Impagidinium torsium* STOVER and HARDENBOL, 1994
Kallo -108.2 m; slide 1. 500x.
- Figures 14, 15. *Impagidinium torsium* STOVER and HARDENBOL, 1994
Kallo -108.2 m; slide 2. 500x.
- Figure 16. *Horologinella* ? sp. 1
Kallo -94 m; slide 1. 1000x.
- Figure 17. *Horologinella* ? sp. 1
Kallo -94 m; slide 1. 1000x.
- Figure 18. *Impletosphaeridium machaeroides* STOVER and HARDENBOL, 1995
Kallo -110.5 m; slide 2. 500x.
- Figure 19. *Gerlachidium* ? sp. indet.
Kallo -107 m; slide 2. 500x.
- Figure 20 : *Gerlachidium aechmophorum* (BENEDEK, 1972)
Kallo -98 m; slide 3. 500x.

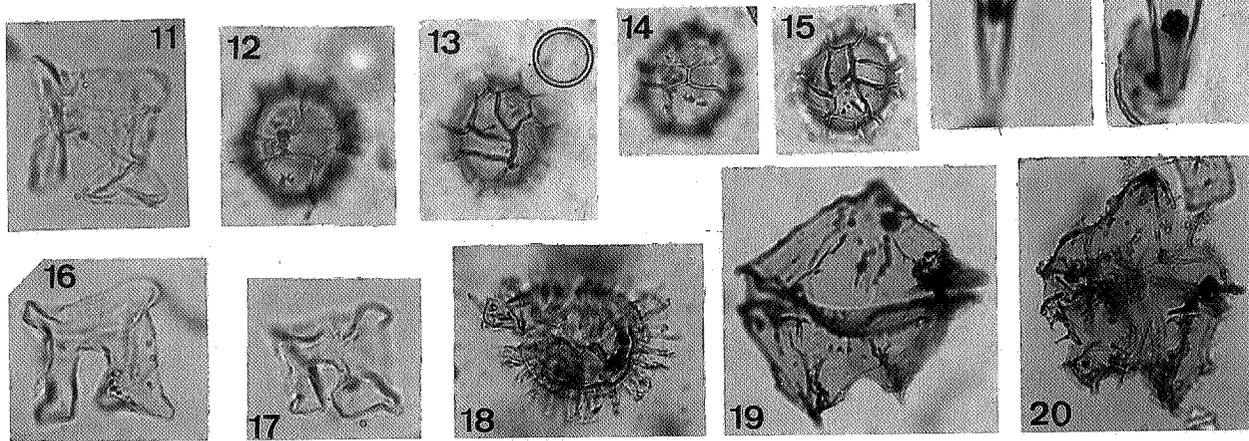
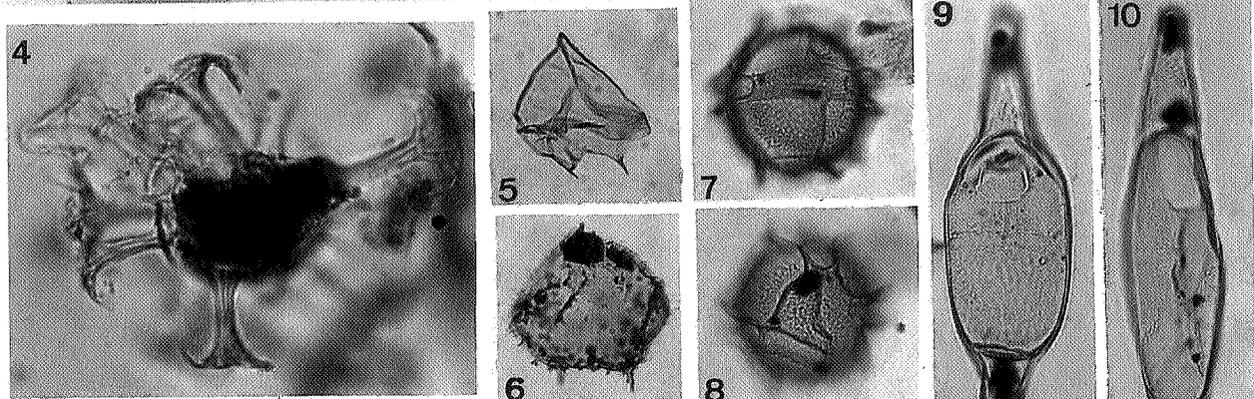
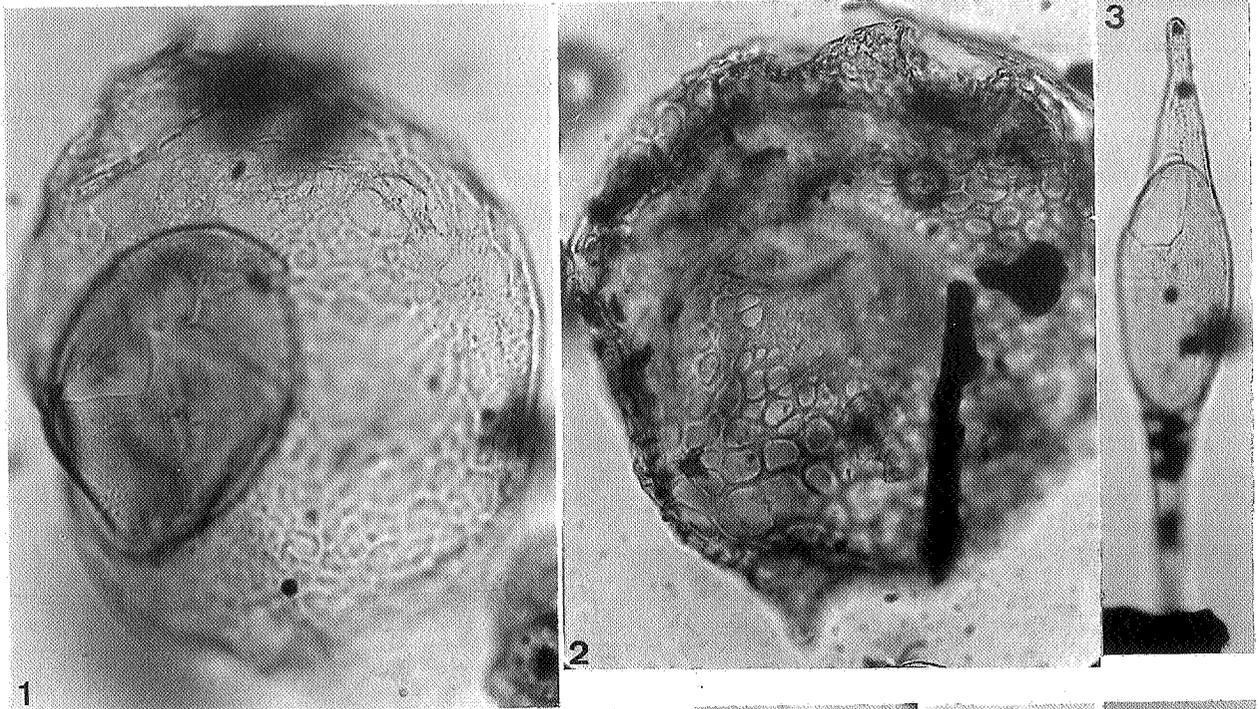
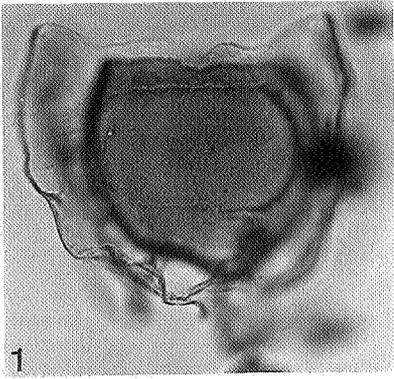
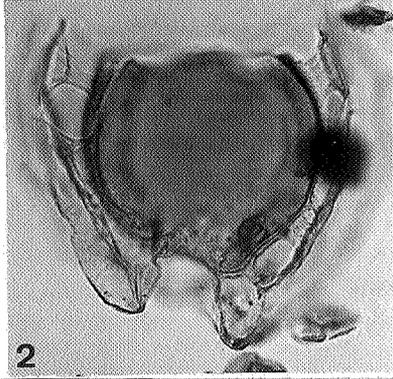


PLATE 3

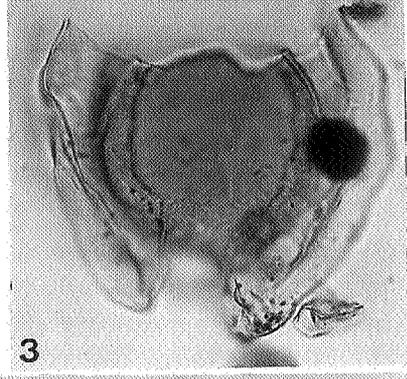
- Figures 1, 2, 3. *Membranophoridium aspinatum* GERLACH, 1961 forma B
Kallo -107 m; slide 1. 500x.
- Figures 4, 5. *Membranophoridium aspinatum* GERLACH, 1961 forma B
Kallo -108 m; slide 3. 500x.
- Figures 6, 7. *Membranophoridium aspinatum* GERLACH, 1961 forma B
Kallo -107 m; slide 2. 500x.
- Figures 8, 9. *Membranophoridium aspinatum* GERLACH, 1961 forma B
Kallo -106 m; slide 1. 500x.
- Figures 10, 11. *Membranophoridium aspinatum* GERLACH, 1961 forma B
Kallo -106 m; slide 1. 500x.
- Figure 12. *Elytrocysta breva* STOVER and HARDENBOL, 1994
Kallo -94 m; slide 1. 500x.
- Figure 13. *Elytrocysta breva* STOVER and HARDENBOL, 1994
Kallo -100 m; slide 2. 500x.
- Figure 14. *Elytrocysta breva* STOVER and HARDENBOL, 1994
Kallo -92.6 m; slide 1. 500x.
- Figure 15. *Elytrocysta breva* STOVER and HARDENBOL, 1994
Kallo -94 m; slide 1. 500x.



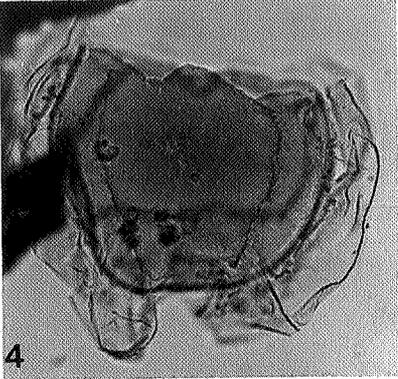
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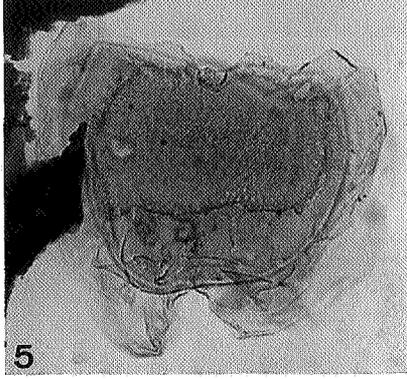
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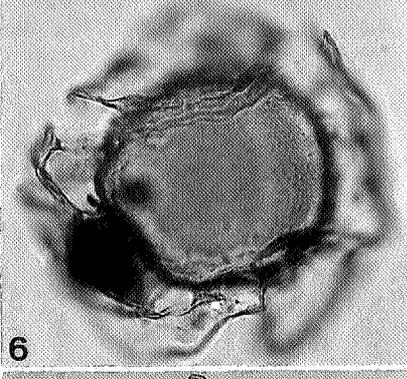
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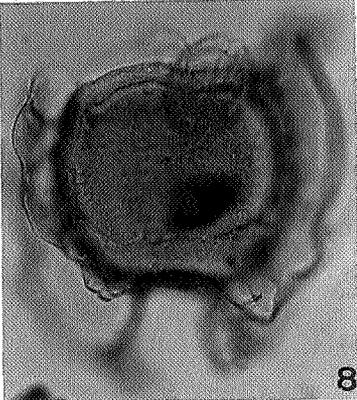
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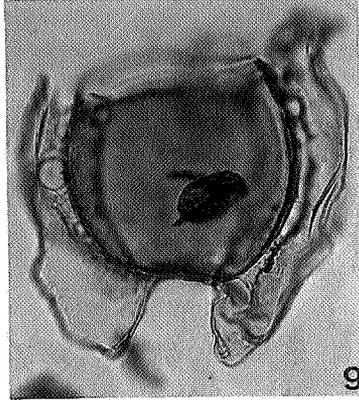
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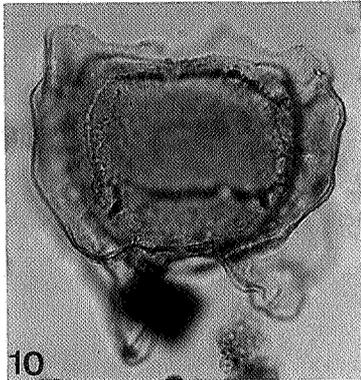
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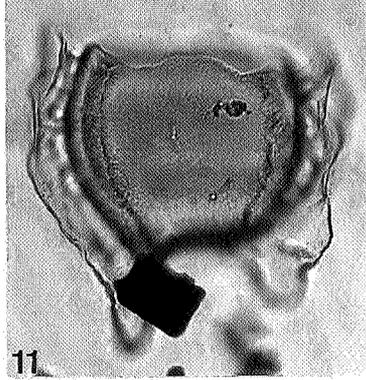
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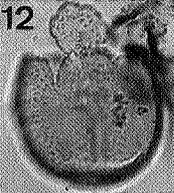
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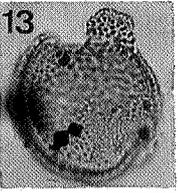
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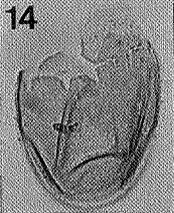
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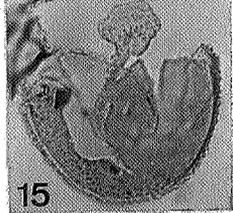
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PLATE 4

- Figure 1. *Hystrichokolpoma* aff. *cinctum* KLUMPP, 1953
Kallo -107 m; slide 3. 500x.
- Figures 2, 3. *Hystrichokolpoma* aff. *cinctum* KLUMPP, 1953
Kallo -106 m; slide 1. 500x.
- Figures 4, 5. *Gerdiocysta conopeum* LIENGJARERN, COSTA and DOWNIE, 1980
Kallo -98 m; slide 4. 500x.
- Figures 6, 7. *Gerdiocysta conopeum* LIENGJARERN, COSTA and DOWNIE, 1980
Kallo -96 m; slide 1. 500x.
- Figure 8. *Vozzhennikovia spinula* STOVER and HARDENBOL, 1994
Kallo -96 m; slide 2. 500x.
- Figure 9. *Selenopemphix selenoides* BENEDEK, 1972
Kallo -106 m; slide 1. 500x.
- Figure 10. *Selenopemphix selenoides* BENEDEK, 1972
Kallo -96 m; slide 1. 500x.
- Figures 11, 12. *Areoligera semicirculata* (MORGENROTH, 1966)
Kallo -106 m; slide 1. 500x.
- Figure 13. *Selenopemphix* sp. aff. *S. nephroides* BENEDEK, 1972 - *S. selenoides* BENEDEK, 1972.
Kallo -108 m; slide 1. 500x.

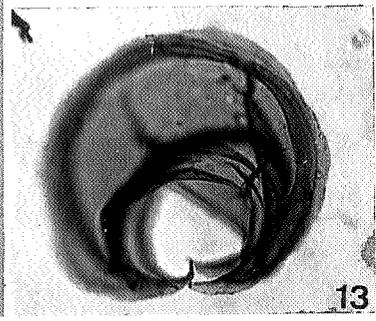
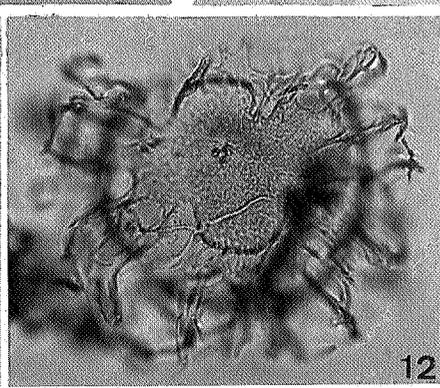
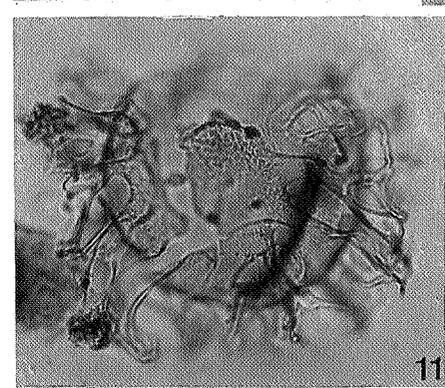
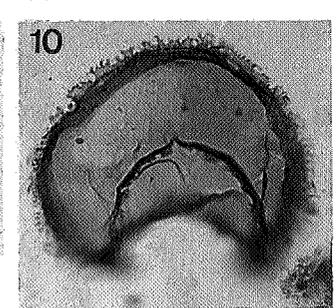
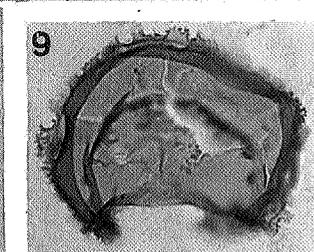
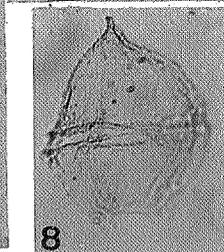
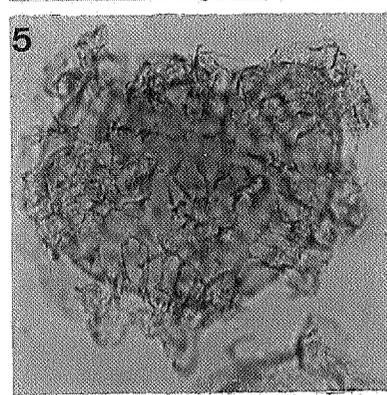
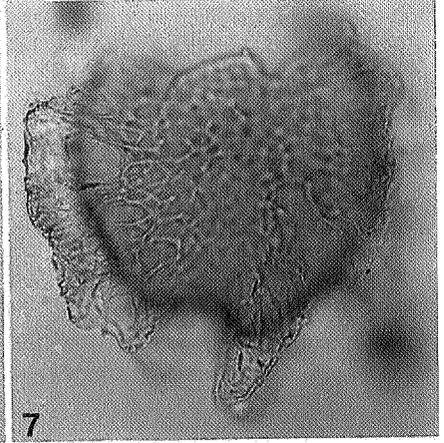
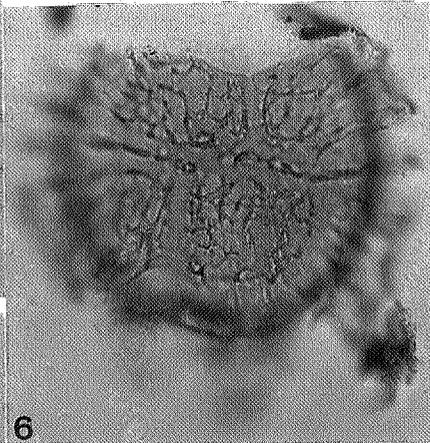
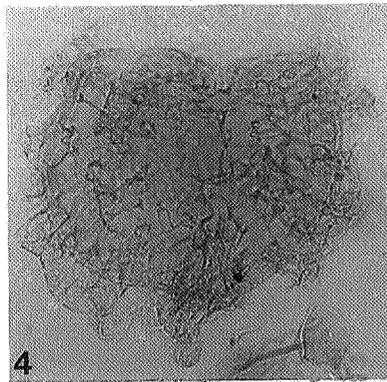
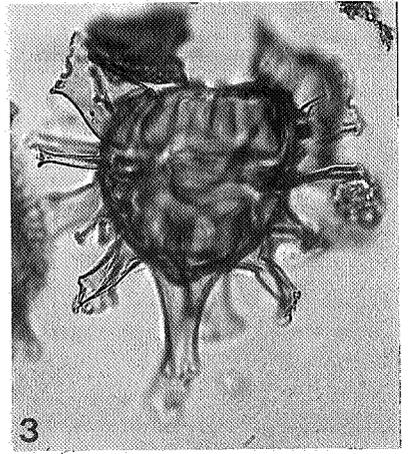
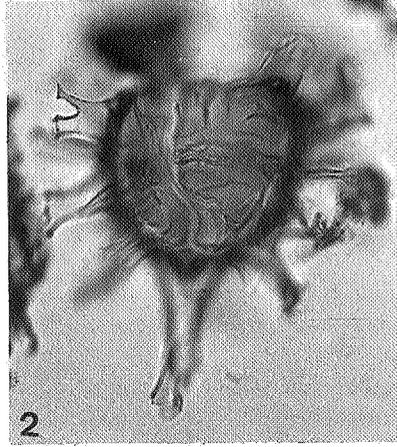
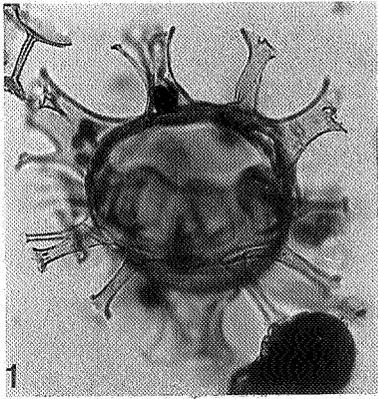


PLATE 5

- Figures 1, 2. ? *Areoligera semicirculata* (MORGENROTH, 1966)
Kallo -111 m; slide 3. 500x.
- Figure 3. *Tuberculodinium vancampoe* (ROSSIGNOL, 1962)
Kallo -100 m; slide 2. 500x.
- Figure 4. *Tuberculodinium vancampoe* (ROSSIGNOL, 1962)
Kallo -92.6 m; slide 2. 500x.
- Figure 5. *Caligodinium amiculum* DRUGG, 1970
Kallo -97 m; slide 1. 500x.
- Figure 6. *Wetzeliiella symmetrica incisa* GERLACH, 1961
Kallo -96 m; slide 1. 500x.
- Figure 7. *Vozzhennikovia cearaichia* STOVER and HARDENBOL, 1974
Kallo -100 m; slide 1. 500x.
- Figure 8. *Trigonopyxidia fiscellata* DE CONINCK, 1986
Kallo -108 m; slide 3. 500x.
- Figure 9. *Trigonopyxidia fiscellata* DE CONINCK, 1986
Kallo -108 m; slide 4. 500x.
- Figure 10. *Reticulosphaera* ? sp. A in DE CONINCK, 1995
Kallo -97 m; slide 1. 1000x.
- Figure 11. *Wetzeliiella symmetrica symmetrica* WEILER, 1956
Kallo -106 m; slide 2. 500x.

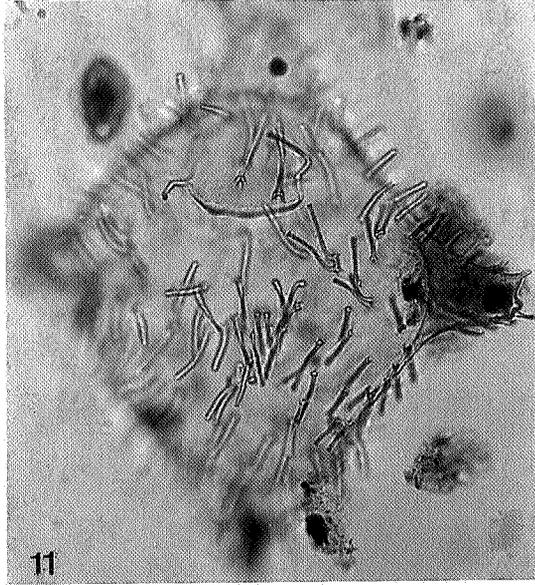
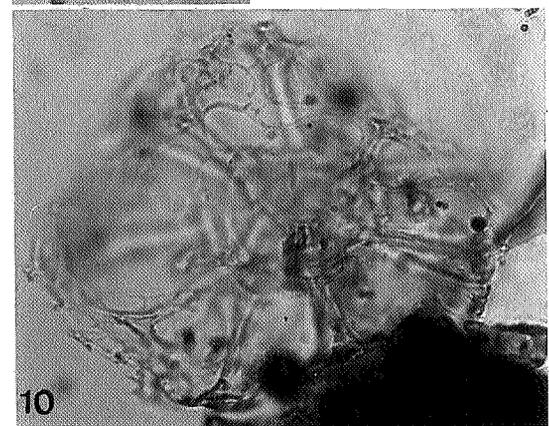
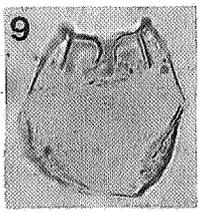
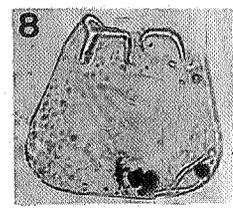
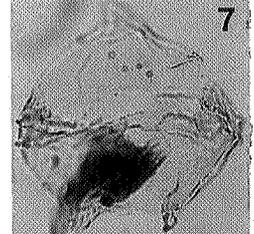
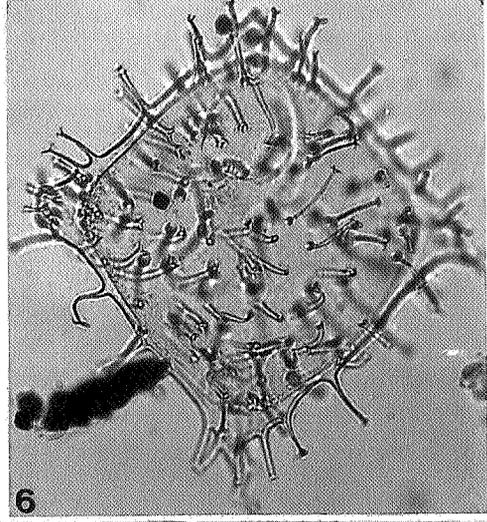
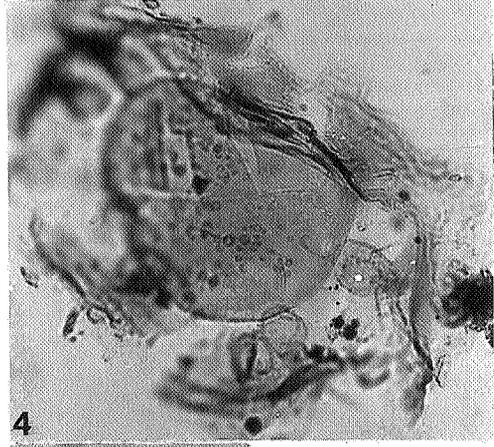
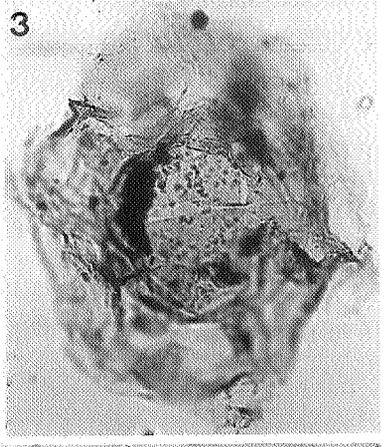
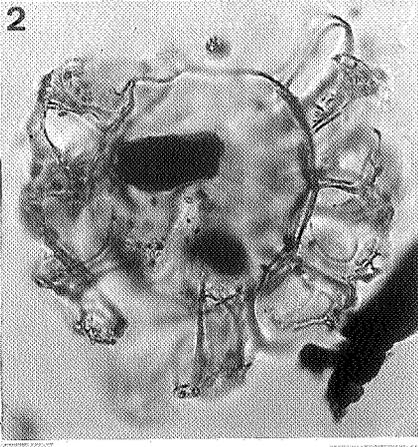
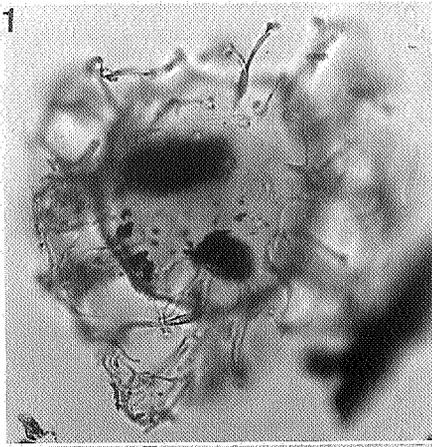


PLATE 6

- Figures 1, 2, 3. *Membranophoridium aspinatum* GERLACH, 1961 forma C
Kallo -100 m; slide 1. 500x.
- Figures 4, 5. *Membranophoridium aspinatum* GERLACH, 1961 forma C
Kallo -100 m; slide 2. 500x.
- Figures 6, 7. *Membranophoridium aspinatum* GERLACH, 1961 forma D
Kallo -97 m; slide 3. 500x.
- Figure 8. *Membranophoridium aspinatum* GERLACH, 1961 forma D
Kallo -96 m; slide 2. 500x.
- Figures 9, 10. *Membranophoridium aspinatum* GERLACH, 1961 forma D
Kallo -94 m; slide 1. 500x.

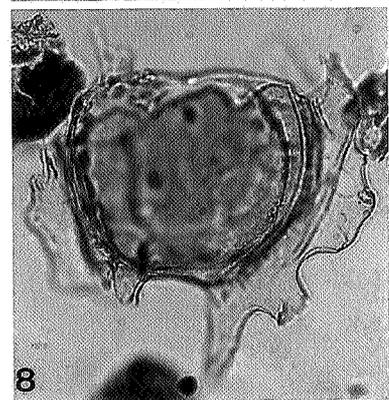
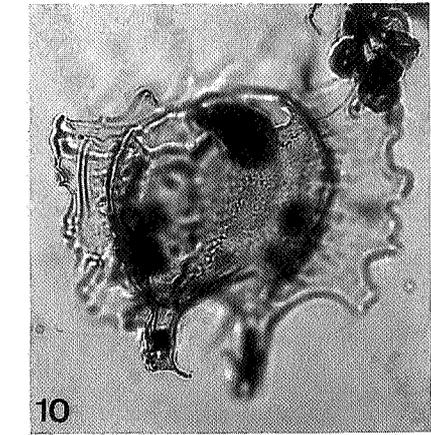
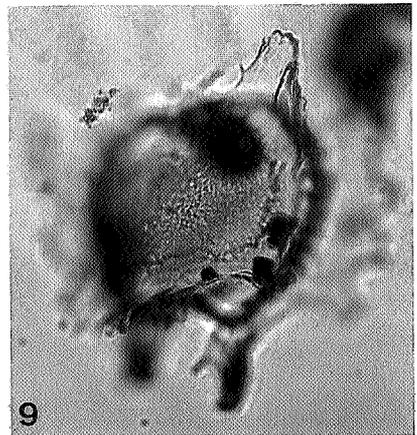
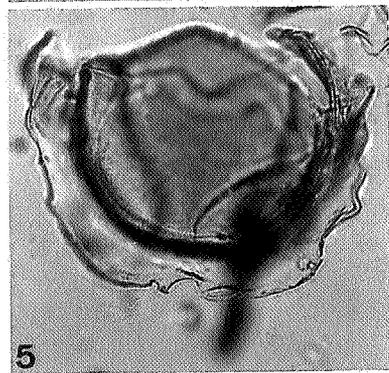
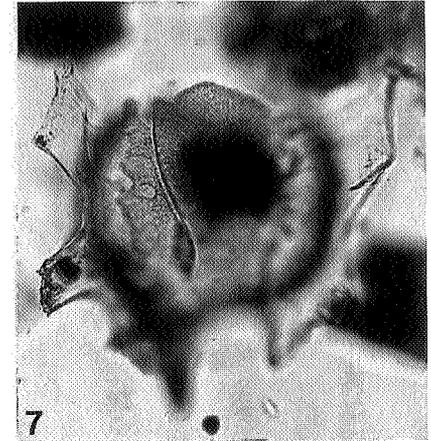
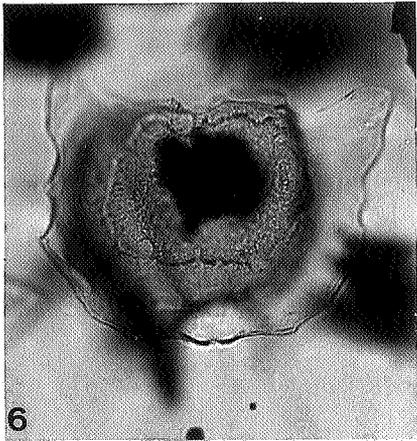
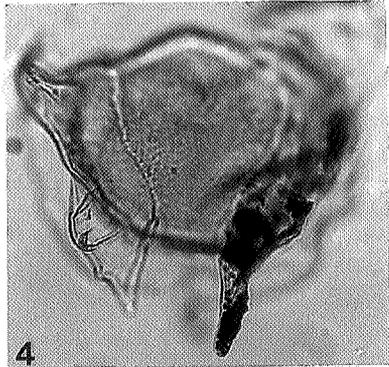
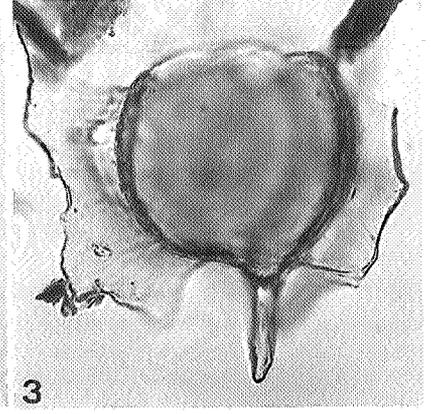
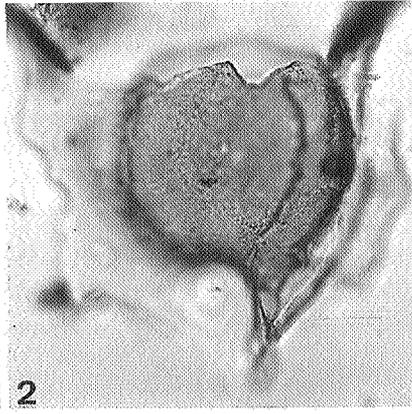
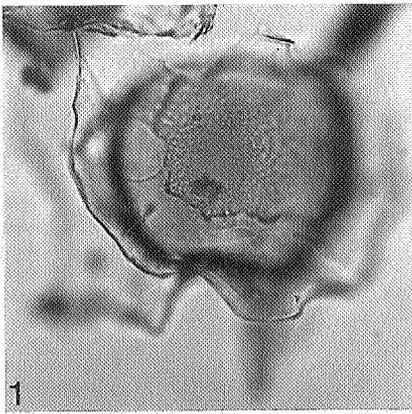
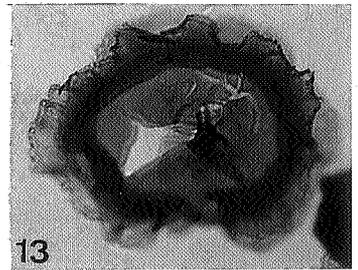
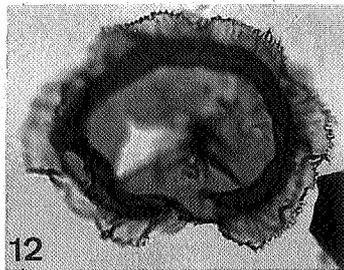
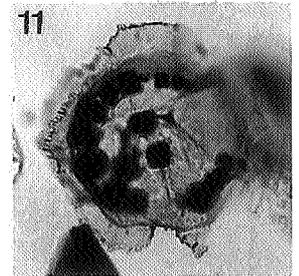
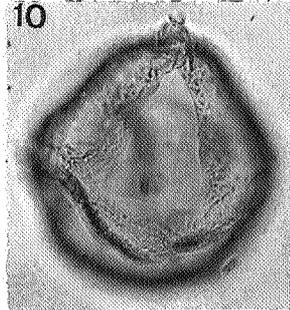
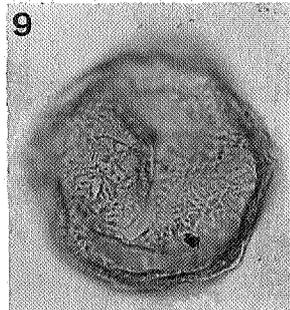
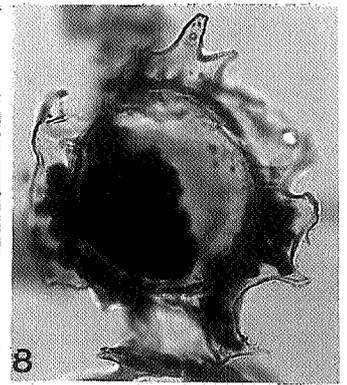
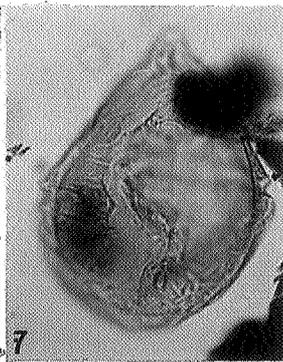
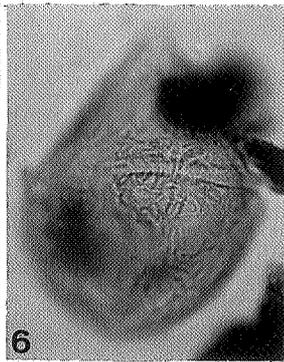
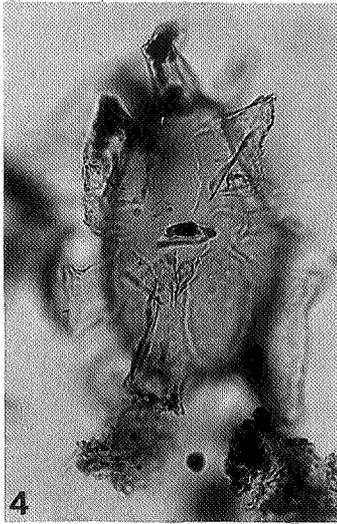
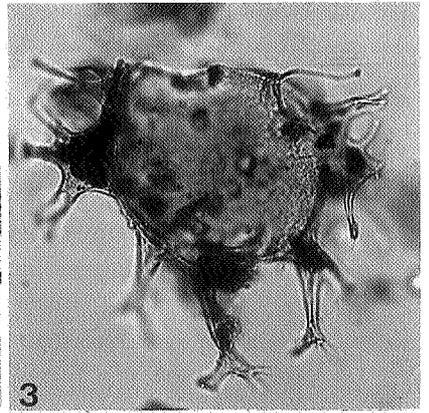
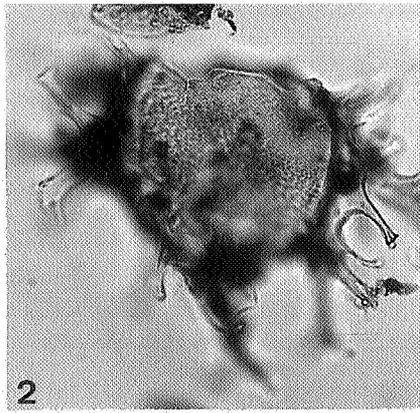
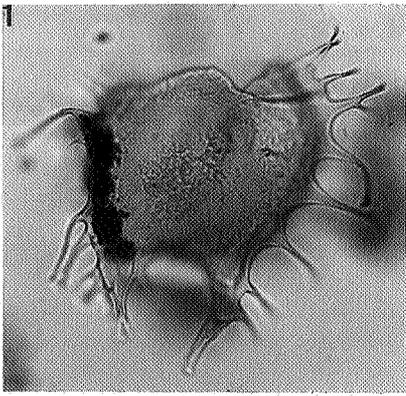


PLATE 7

- Figure 1. *Chiropteridium* sp. indet.
Kallo -94 m; slide 3. 500x.
- Figures 2, 3. *Chiropteridium* sp. indet.
Kallo -94 m; slide 3. 500x.
- Figures 4, 5. *Achilleodinium* aff. *biformoides* (EISENACK, 1954)
Kallo -94 m; slide 3. 500x.
- Figures 6, 7. *Apteodinium spiridoides* BENEDEK, 1972
Kallo -94 m; slide 1. 500x.
- Figure 8. *Membranophoridium aspinatum* GERLACH, 1961 forma D
Kallo -96 m; slide 3. 500x.
- Figures 9, 10. *Apteodinium spiridoides* BENEDEK, 1972
Kallo -94 m; slide 1. 500x.
- Figure 11. *Selenopemphix* ? sp. 3 in HEAD and NORRIS, 1989
Kallo -92.6 m; slide 2. 500x.
- Figures 12, 13. *Selenopemphix* ? sp. 3 in HEAD and NORRIS, 1989
Kallo -92.6 m; slide 2. 500x.



115	112.1	111	110.5	110	109.5	108.2	108	107	106	100	98	97	96	94	92.6
cf. <i>Platycystidia</i> sp. II in MANUM 1976															
<i>Polysphaeridium zoharyi</i> (ROSSIGNOL 1962)															
<i>Prolixosphaeridium granulosum</i> (DEFLANDRE 1937)															
<i>Pseudocecratium pelliferum</i> GOCHT 1957															
<i>Pterodinium</i> ? <i>cingulatum</i> (O. WETZEL 1933)															
<i>Raetia</i> <i>indium</i> <i>evitigrata</i> KIRSCH 1991															
<i>Reticulatosphaera</i> ? sp. A															
<i>Rhombodinium draco</i> GOCHT 1955															
<i>R. glabrum</i> (COOKSON 1956)															
<i>R. perforatum</i> (JAN DU CHENE and CHATEAUNEUF 1975)															
<i>R. pustulosum</i> CHATEAUNEUF 1980															
<i>Rhynchodiniopsis</i> spp. <i>indet.</i>															
<i>Rotnestia borussica</i> (EISENACK 1954)															
<i>Rotnestia</i> sp. <i>indet.</i>															
<i>Samlandia chlamydrophora</i> EISENACK 1954															
<i>Scriniodinium campanulum</i> GOCHT 1959															
<i>Selenopemphix armata</i> BUJAK 1980															
<i>Selenopemphix</i> ? sp. 3 in HEAD and NORRIS 1989															
<i>Sentusidinium</i> spp. <i>indet.</i>															
<i>Spintiferella cornuta</i> (GERIACH 1961)															
<i>Spintiferites pseudofurcatus</i> (KLUMPP 1953)															
S. sp. cf. S. sp. A in POWELL 1986															
<i>Spintiferites</i> spp. <i>indet.</i>															
<i>Stiphrosphaeridium</i> ? spp. <i>indet.</i>															
<i>Svalbardella</i> sp. <i>indet.</i>															
<i>Systematophora placacantha</i> (DEFLANDRE and COOKSON 1955)															
? <i>S. palmita</i> DAVEY 1982															
<i>Tanyosphaeridium</i> spp. <i>indet.</i>															
<i>Tectatodinium pellitum</i> WALL 1967															
T. ? sp. cf. <i>Filipsphaera</i> sp. A in HEAD, NORRIS and MUDIE 1989															
<i>Thalassiphora delicata</i> WILLIAMS and DOWNIE 1966															
<i>T. fenestrata</i> LIENGGJAREN, COSTA and DOWNIE 1980															
<i>T. patula</i> (WILLIAMS and DOWNIE 1966)															
T. ? cf. <i>pansa</i> STOVER 1977															

	115	112.1	111	110.5	110	109.5	108.2	108	107	106	100	98	97	96	94	92.6
M. stellatum DEFLANDRE 1942																
M. ? tubulispinosum DE CONINCK 1976																
M. wattonense WALL 1965																
Pacillobimorpha spinosa (COOKSON 1965)	:															
P. triradialata DE CONINCK 1986	.															
Quadrina pallida BUJAK 1980	.															
Verhachium europaeum STOCKMANS and WILLIÈRE 1960	I															
Verhachium ? - Micrhystriidum ? spp. indet.			I													
Number of dinocyst species reworked from the Jurassic	1	1	3	4	4	2	0	0	4	4	1	0	6	3	3	5
Cretaceous	1	1	7	8	6	12	2	1	5	6	6	5	12	9	8	5
Paleocene-Lutetian	0	0	2	3	0	1	2	3	5	3	2	2	1	2	1	0
Number of contemporaneous dinocyst species	74	54	55	68	61	66	63	69	63	66	55	69	47	57	57	64

TABLE 3 :

Reworked dinocyst taxa

Kallo borehole: depth in meters below the surface	115	112.1	111	110.5	110	109.5	108.2	108	107	106	100	98	97	96	94	92.6
<u>Jurassic</u>																
<i>Acanthaulax</i> cf. <i>aculeata</i>						.										
<i>Criboperidinium</i> spp. indet.			I	.	.	.
<i>Ctenidodinium</i> sp. indet.													.			?
<i>Dingodinium tuberosum</i>		
<i>Glossodinium dimorphum</i>														.		
<i>Hystrichogonyaulax cladophora</i>									.							
<i>Leptodinium</i> spp. indet.								
<i>Lithodinia</i> ? sp. indet.													.			
<i>Netrelytron</i> ? sp. indet.																.
<i>Pareodinia</i> ? sp. indet.			.						.							
<i>Sentusidinium</i> spp. indet.		
<i>Valensiella</i> spp. indet.				.												
<i>Xenicodinium densispinosum</i>													?			
<u>Cretaceous</u>																
<i>Apteodinium granulatum</i>					
<i>Callaiosphaeridium asymmetricum</i>						.					.					
<i>Canninginopsis</i> sp. indet.														.		
<i>Chlamydophorella</i> spp. indet.		
<i>Circulodinium colliveri</i>															.	.
<i>Cyclonephelium distinctum</i>							
<i>Cyclonephelium compactum</i>						.								.		
<i>Cyclonephelium hystrix</i>			.			.							.			
<i>Gardodinium</i> sp. indet.						.									.	
<i>Gonyaulacysta fastigiata</i>		
<i>Hystrichodinium pulchrum</i>								?
<i>Hystrichodinium voigtii</i>												
<i>Imbatodinium</i> cf. <i>radiculatum</i>													.			
<i>Kleithriasphaeridium</i> sp. indet.													?		.	
<i>Muderongia</i> sp. indet.													.			
<i>Odontochitina</i> sp. indet.	

Table 3. Distribution in the Kallio section of reworked dinocyst taxa and of *Pediastrum*.

	115	112.1	111	110.5	110	109.5	108.2	108	107	106	100	98	97	96	94	92.6
<i>Oligosphaeridium</i> spp. indet.			.	.		:							.	.		
<i>Phoberocysta neocomica</i>									.	.	.					
<i>Prolixosphaeridium granulosum</i>				.									.			
<i>Pseudoceratium pelliferum</i>		
<i>Pterodinium aliferum</i>					.											
<i>Pterodinium cingulatum</i>				.			.									
<i>Raetiadinium evitigracia</i>			.			.					.		:		.	
<i>Rhynchodiniopsis</i> spp. indet.													.	.	.	
<i>Scriniodinium campanulum</i>							
<i>Stiphrosphaeridium</i> ? spp. indet.			.	:		.			.	?	.					
<i>Systematophora palmula</i>																?
<u>Paleocene - Ypresian - Lutetian</u>																
P-Y <i>Adnatosphaeridium robustum</i>									.							
P-Y <i>Hystrichosphaeridium tubiferum</i>				.			.	.								
P-Y <i>Kallosphaeridium breviparbatum</i>																
P-Y <i>Thalassiphora delicata</i>			.		.											
Y <i>Charlesdownia</i> aff. <i>clathrata</i>										.						
Y <i>Charlesdownia crassoramosa</i>			.										.			
Y <i>Dracodinium varielongitudum</i>								.					.			
Y <i>Eatonicysta ursulae</i>									
Y <i>Membranilarnacia</i> sp. A							?		.	?	.		.	.		
Y-L <i>Lentinia wetzelii</i>			?	.					.	.						
L <i>Wetzeliella pachyderma</i>									.							
<u>Fresh water algae: <i>Pediastrum</i> spp. indet.</u>								II	II				I			

Table 4

Niel boring 43W-270

	Watervliet	Wintham		Ruisbroek	
	Clay	Silt		Sand	
	-30.00	-29.30	-28.5 to -28	-27.80	-27.25
<i>Areosphaeridium diktyoplokus</i>		x	x	x	
<i>Cerebrocysta bartonensis</i>		x	x	x	
<i>Heteraulacacysta porosa</i>	x	x		x	
<i>Membranophoridium aspinatum</i> forma A	x				
<i>Charlesdownia clathrata-coleothrypta</i>	x	x	x		
<i>Glaphyrocysta semitecta</i>			x	x	x
<i>Paucilobimorpha spinosa</i>		x	x	x	
<i>Rhombodinium draco</i>	x				
<i>Rhombodinium perforatum</i>	x	x		x	
<i>Glaphyrocysta</i> aff. <i>inculta</i>	x	x		?	
<i>Thalassiphora fenestrata</i>	x				
<i>Lophocysta</i> sp. indet.	?		x	x	x
<i>Homotryblium caliculum</i>	x				
<i>Rhombodinium glabrum</i>				x	
<i>Ascostomocystis potane</i>	x				x
<i>Planoperidinium gracile</i>	x	x	x	x	x
<i>Operculodinium</i> sp. cf. <i>O. placitum</i>					x
<i>Gerlachidium</i> sp. indet.					x
<i>Selenopemphix selenoides</i>		?			x
<i>Gerdicocysta conopeum</i>					x
<i>Areoligera semicirculata</i>	x				?
<i>Wetzeliella symmetrica</i>					x
<i>Vozzhennikovia cearaichia</i>					x
<i>Reticulatosphaera?</i> sp. A					x
<i>Membranophoridium aspinatum</i> forma C					x
<i>Achilleodinium</i> aff. <i>biformoides</i>			?		?

Table 4. Distribution in the Niel 43W-270 section of dinocyst taxa considered to be biostratigraphically significant.

Table 5

Hingene-Wintham boring 42E-212

	Wintham Silt		Ruisbroek Sand
	-29.95	-27.95	-26.95
<i>Areosphaeridium diktyoplokus</i>	x	x	
<i>Cerebrocysta bartonensis</i>	x		
<i>Heteraulacacysta porosa</i>	x	x	
<i>Membranophoridium aspinatum</i> forma A		x	
<i>Charlesdowniea clathrata</i>	x	x	
<i>Glaphyrocysta semitecta</i>	x	x	x
<i>Rhombodinium perforatum</i>	x	x	
<i>Glaphyrocysta</i> aff. <i>inculta</i>		x	
<i>Lophocysta</i> sp. indet.		x	x
<i>Thalassiphora fenestrata</i>	x	x	
<i>Rhombodinium glabrum</i>		x	
<i>Ascostomocystis potane</i>			?
<i>Planoperidinium gracile</i>			x
<i>Selenopemphix selenoides</i>			x
<i>Gerdiocysta conopeum</i>			x
<i>Operculodinium</i> sp. cf. <i>O. placitum</i>			x
<i>Areoligera semicirculata</i>	x		
<i>Svallardella</i> sp. indet.			?
<i>Wetziella symmetrica</i>			x
<i>Vozzhennikovia cearaichia</i>			x
<i>Reticulosphaera?</i> sp. A	?		x
<i>Membranophoridium aspinatum</i> forma C			x
<i>Achilleodinium</i> aff. <i>biformoides</i>			x
<i>Chiropteridium</i> sp.			x

Table 5. Distribution in the Hingene-Wintham 42E-212 section of dinocyst taxa considered to be biostratigraphically significant.

Table 6

Terhagen boring 58W-213

	Watervliet Clay	Wintham Silt
	-38.50	-37.55
<i>Areosphaeridium diktyoplokus</i>	x	x
<i>Heteraulacacysta porosa</i>		x
<i>Membranophoridium aspinatum</i> forma A	x	
<i>Charlesdownia clathrata-coleothrypta</i>	x	x
<i>Galphyrocysta semitecta</i>	x	
<i>Paucilobimorpha spinosa</i>		x
<i>Rhombodinium draco</i>	x	x
<i>Rhombodinium perforatum</i>	x	x
<i>Glaphyrocysta</i> aff. <i>inculta</i>	x	x
<i>Lophocysta</i> sp. indet.	x	
<i>Thalassiphora fenestrata</i>	x	x
<i>Thalassiphora reticulata</i>		x

Table 6. Distribution in the Terhagen 58W-213 section of dinocyst taxa considered to be biostratigraphically significant.